

Does nitrogen availability have greater control over the formation of tropical heath forests than water stress? A hypothesis based on nitrogen isotope ratios

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ABSTRACT

Global scale analyses of soil and foliage $\delta^{15}\text{N}$ have found positive relationships between $\delta^{15}\text{N}$ and ecosystem N loss (suggesting an open N cycle) and a negative relationship between $\delta^{15}\text{N}$ and water availability. We show here that soils and leaves from tropical heath forests are depleted in ^{15}N relative to 'typical' forests suggesting that they have a tight N cycle and are therefore limited by N rather than by, often suggested, water availability.

KEYWORDS: acid soils, kerangas, nitrous oxide, stable isotopes

Tem a disponibilidade de nitrogênio maior controle sobre a formação de florestas tropicais de campinarana do que o estresse hídrico? Uma hipótese fundamentada nas razões isotópicas do nitrogênio

RESUMO:

Análises de $\delta^{15}\text{N}$ solo e foliar, em escala global, encontraram relações positivas entre $\delta^{15}\text{N}$ e a perda de N do ecossistema (sugerindo um ciclo aberto de N), e uma relação negativa entre o $\delta^{15}\text{N}$ e a disponibilidade de água. Mostramos neste trabalho que os solos e a vegetação de florestas de campinarana são empobrecidas em ^{15}N em relação a florestas 'típicas', sugerindo que elas têm um ciclo fechado de N e são, desta forma, limitadas pelo N, ao invés da disponibilidade de água, como frequentemente sugerido.

PALAVRAS-CHAVE: solos ácidos, campinarana, óxido nitroso, isótopos estáveis.

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Determining the factors that lead to differing species distributions in tropical forests, especially in relation to soil type, has been an important question in tropical ecology for many years and still remains a contemporary problem (Fine *et al.* 2004; Russo *et al.* 2005). Tropical heath forests (HF) (also known as white sand forests in the Neotropics) are a formation that are very different to 'typical' lowland evergreen rain forests (LERFs) having more acidic soil, lower stature, a more even canopy, thicker smaller leaves and a greater proportion of trees in smaller size classes (Brünig 1974; Proctor 1999). There are a number of lines of evidence that suggest HFs have lower nitrogen (N) availability than LERFs. For example, Turner *et al.* (2000) found lower foliar N in a range of species from HFs in Brunei and Moran *et al.* (2000) found lower N in soil solutions and a reduction in N returned *via* litterfall fluxes in the same forests. In addition, Fine *et al.* (2005) found lower mineral N availability in Peruvian HFs. In contrast, Brünig (1971, 1974) considers that heath forests are formed due to water limitation which may be especially prevalent on sandy soils.

Examination of N stable isotopes is a powerful methodology that can assess N inputs, outputs and transformations in an ecosystem (Martinelli *et al.* 1999; Robinson 2001; Amundsen *et al.* 2003); N availability to plants and trees can be measured by determining stable N isotope abundance. As an example, tropical rain forest soils and leaves are ¹⁵N enriched compared to temperate forests (Martinelli *et al.* 1999). The reason for this difference is related to the nature of the greater loss of N from tropical forests: there is preferential leaching and gaseous loss of the lighter ¹⁴N isotope thereby leaving the soils enriched in ¹⁵N. However, another factor which leads to changes in δ¹⁵N between soils is a change in moisture regime as a decrease in δ¹⁵N with increasing rainfall has been found (Austin and Vitousek 1998; Handley *et al.* 1999; Schuur and Matson 2001). This is hypothesised to be because of a changing preference for organic forms of N in wetter soils (Handley *et al.* 1999).

The aim of this study was to compare δ¹⁵N values of soil and plant material from tropical HFs and LERFs and test the null hypothesis of no differences in δ¹⁵N signatures and hence minimal differences in N cycling between the two forest types.

Soil and leaf samples were taken from HF and LERF areas of six tropical sites (Table 1). From four of these sites soils (0-10 cm) were collected and from two of the sites leaves of the sub-canopy species *Protium subseratum* Engl. (Burseraceae) were collected; this species was chosen as it is one of the only species which is common in both HFs and LERFs (Fine *et al.* 2005).

Samples were analysed for δ¹⁵N using a ThermoFinnigan Delta^{plus} IRMS interfaced with a CE Instruments 1112 Flash elemental analyser via a ConFlo III (or a Europa 20-20 IRMS

Table 1 - Locations and rainfall of sites from where soils and leaves were collected for δ¹⁵N analysis.

Site	Latitude, Longitude	Rainfall
Soils		
Bako, Sarawak, Malaysia	1°43' N, 110°27' E	4048 mm (Harrison 2004 for Kuching)
Barito Ulu, Kalimantan, Indonesia	0°06' S, 114°0' E	3800 mm (Brearley <i>et al.</i> 2007)
Kabili-Sepilok, Sabah, Malaysia	5°52' N, 117°56' E	3000 mm (Brearley <i>et al.</i> 2003)
Mabura Hill, Guyana	5°13' N, 58°13' W	2700 mm (Perreijn 2002)
Leaves		
Iquitos, Peru	3°54' S, 73°31' W	3087 mm (Marengo 1998)
Rio Blanco, Peru	6°05' S, 73°42' W	3087 mm (Marengo 1998)

Where LERF and HF sites are separated (e.g. Peru), the location is the mid-point between the two sites.

interfaced with an ANCA-GSL elemental analyser for the Guyana samples). N isotope ratios are expressed as delta (δ) notation which is the per mille deviation from atmospheric N: δ¹⁵N (‰) = [(R_{sample}/R_{standard}) - 1] x 1000 where R is ¹⁵N/¹⁴N. Precision (standard deviation) of duplicated measurement of sixteen samples on the Delta^{plus} IRMS was 0.38 ‰.

Soil δ¹⁵N values were significantly lower in the HFs when compared with LERF ($F_{1,89} = 190, p < 0.001$; Figure 1a). Leaf δ¹⁵N values of *Protium subseratum* were also significantly lower in the HF when compared with LERF ($F_{1,20} = 118, p < 0.001$; Figure 1b). There was a negative correlation between rainfall at each site and soil δ¹⁵N ($r = -0.872$) although this was not statistically significant ($p = 0.128$).

Numerous studies have shown that a climatic shift of increasing rain fall leads to a decrease in δ¹⁵N of soil and/or foliage (Austin and Vitousek 1998; Handley *et al.* 1999;

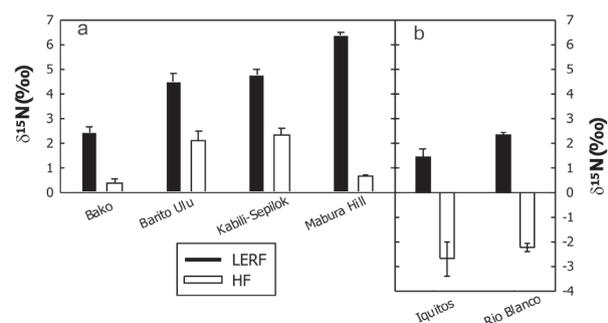


Figure 1 - Nitrogen isotope (δ¹⁵N) values of soils from four tropical sites (a) and leaves (b) from two tropical sites with lowland evergreen rain forests (LERF) and heath forests (HF). Soil type ($F_{1,89} = 190, p < 0.001$), site ($F_{3,89} = 13.8, p < 0.001$) and the soil type x site interaction ($F_{3,89} = 32.7, p < 0.001$) all had a significant effect on the soil sample δ¹⁵N values. Soil type ($F_{1,20} = 118, p < 0.001$), but neither site ($F_{1,20} = 2.74, p = 0.11$) nor the soil type x site interaction ($F_{1,20} = 0.31, p = 0.58$) had a significant effect on the leaf δ¹⁵N values.

Schuur and Matson 2001; Amundsen *et al.* 2003; Sah *et al.* 2006) both in temperate and tropical sites. Following this pattern, water limitation would be expected to lead to increases in $\delta^{15}\text{N}$. If, therefore, tropical HF soils were water limited over LERF soils we would expect an increase in $\delta^{15}\text{N}$, not the decrease which was actually seen. The change in $\delta^{15}\text{N}$ values are most likely to result from an altered N cycle in the HFs and almost certainly from reduced losses of ^{15}N depleted compounds.

Whilst there are often differences in available soil N between forest types there has been little evidence for differences in mineralisation rates. Can our isotopic results therefore be reconciled with those of Vernimmen *et al.* (2007) and Fine *et al.* (unpublished data) who found minimal differences in N mineralisation between HF and LERF soils at the Barito Ulu and Iquitos sites respectively? Yes, in so far that for isotopic fractionation to occur, there must be losses of ^{15}N depleted compounds from soils. This can occur through nitrate leaching or denitrification to nitrous oxide. Indeed, we know that emissions of nitrous oxide can be very high from tropical forest soils (Houlton *et al.* 2006). At our Guyana study site, Perreijn (2002) found that denitrification rates were lower in HF than LERF and we therefore hypothesised that this, more generally, is the cause of the differences in soil $\delta^{15}\text{N}$ between the two forest types as suggested by Houlton *et al.* (2006) to explain differences in soil $\delta^{15}\text{N}$ in relation to climate.

This paper therefore suggests that nitrogen availability is more important in controlling the formation of HFs than water availability confirming the results of previous investigators who found little evidence for drought tolerance among HF trees (Becker 1996; Nagy and Proctor 2000; Sobrado 2010). Nagy and Proctor (2000) also showed that foliar $\delta^{13}\text{C}$ values were not different between canopy trees in HF and LERF at Barito Ulu adding further support to our hypothesis that water limitation is not of over-riding importance in leading to the HF formation. We suggest that further examination of N cycling patterns in both forest types are continued to examine exactly where in this complex cycle differences between the forest types occur. However, whilst ruling out the over-riding importance of water stress it does not rule out that N cycling may be limited by extreme soil acidity as suggested by Proctor (1999).

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